

Projections of precipitation extremes, tropical cyclones, extratropical cyclones and blockings

Akio KITO

Meteorological Research Institute, Tsukuba, Japan

- *application of high-resolution model for future climate projections*
- *multi-decades integrations needed*
- poster by H. Murakami tomorrow on hurricane changes

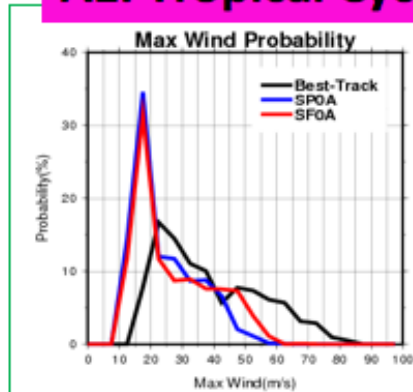


Projection of the Change in Weather Extremes Using Super-High-Resolution Atmospheric Models in the KAKUSHIN Program

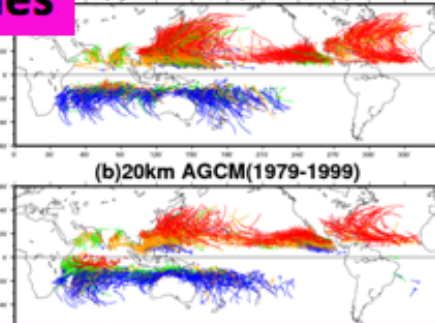


Akio Kitoh (MRI/JMA), Shoji Kusunoki (MRI/JMA), Eiichi Nakakita (DPRI/Kyoto-Univ.),
Kunivoshi Takeuchi (ICHARM/PWRI) and Hiroki Kondo (AESTO) with the project team

A1. Tropical Cyclones

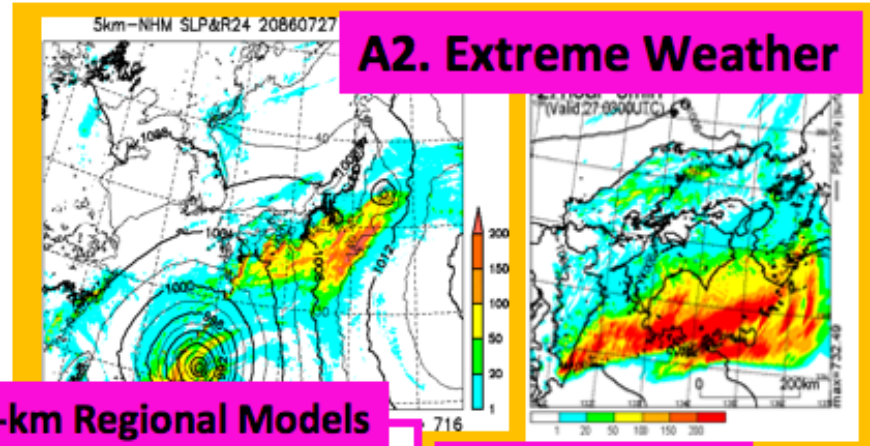


(a) Best-Track (1979-1999)



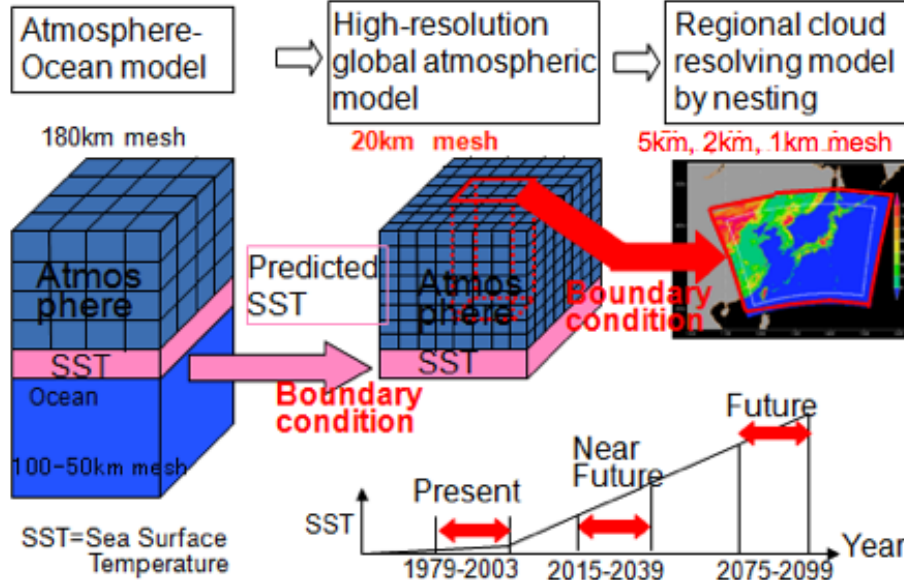
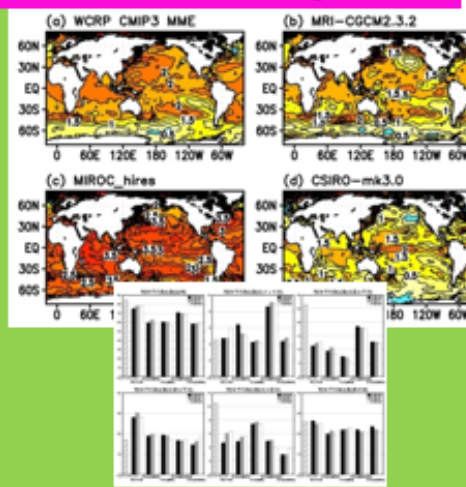
(b) 20km AGCM (1979-1999)

A2. Extreme Weather

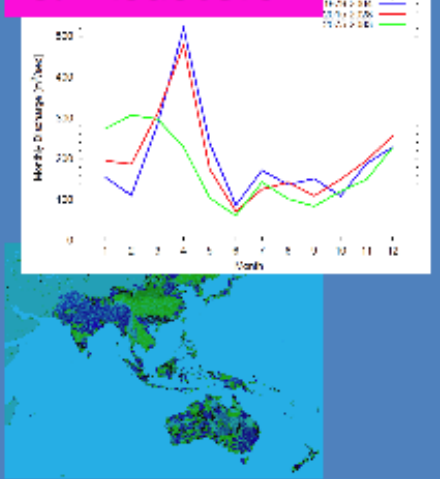


20km Global and 1-, 2- & 5-km Regional Models

B. Uncertainty

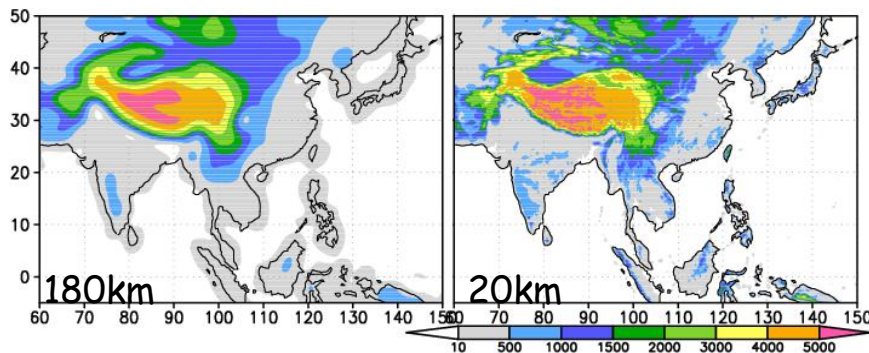


C. Flood & Disasters

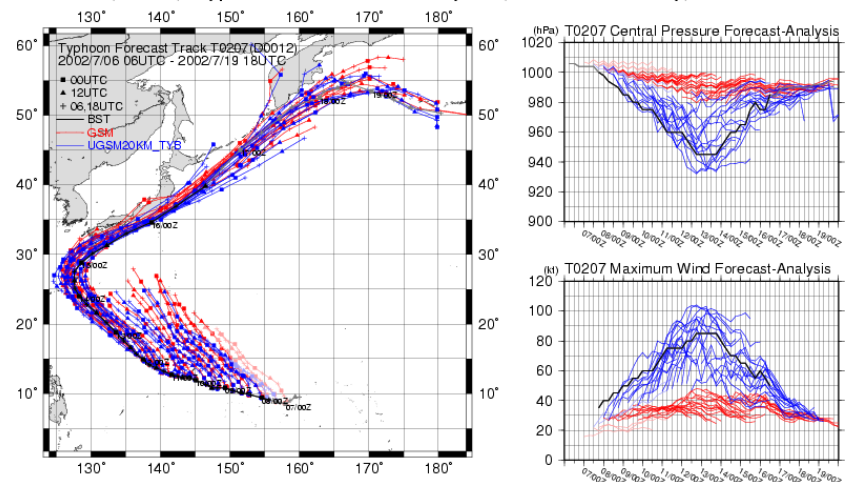


MRI/JMA Atmospheric GCM

- JMA : **Operational global NWP model from Nov 2007**
- MRI : Next generation climate model
- Based on operational JMA-GSM
- Resolution: **TL959 (20km) with 60 layers (0.4 hPa)**
- Time integration: **Semi-Lagrangian Scheme** (Yoshimura, 2004)
- Physics
 - SW radiation: Shibata & Uchiyama (1992)
 - LW radiation: Shibata & Aoki (1989)
 - Cumulus convection: Prognostic Arakawa-Schubert (Randall and Pan, 1993)
 - Land hydrology: MJ-SiB: SiB with 4 soil-layers and 3 snow-layers
 - Clouds: large-scale condensation, Cumulus, stratocumulus
 - PBL: Mellor & Yamada (1974,1982) level-2 closure model
 - Gravity wave drag: Iwasaki et al. (1989) + Rayleigh friction

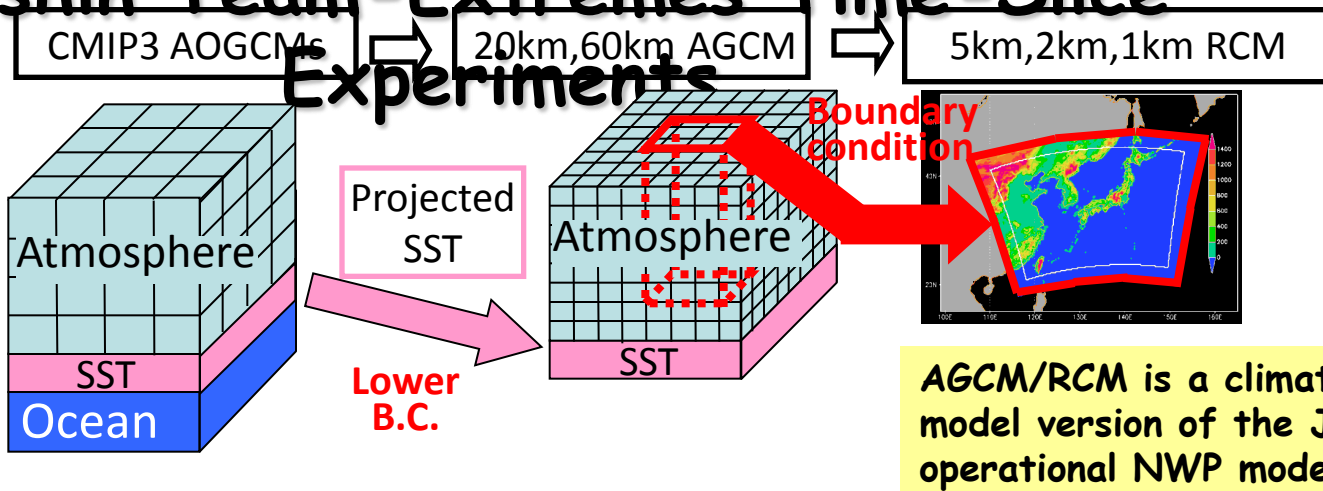


T0207(D0012) Typhoon Forecast and Analysis (Track and Intensity)

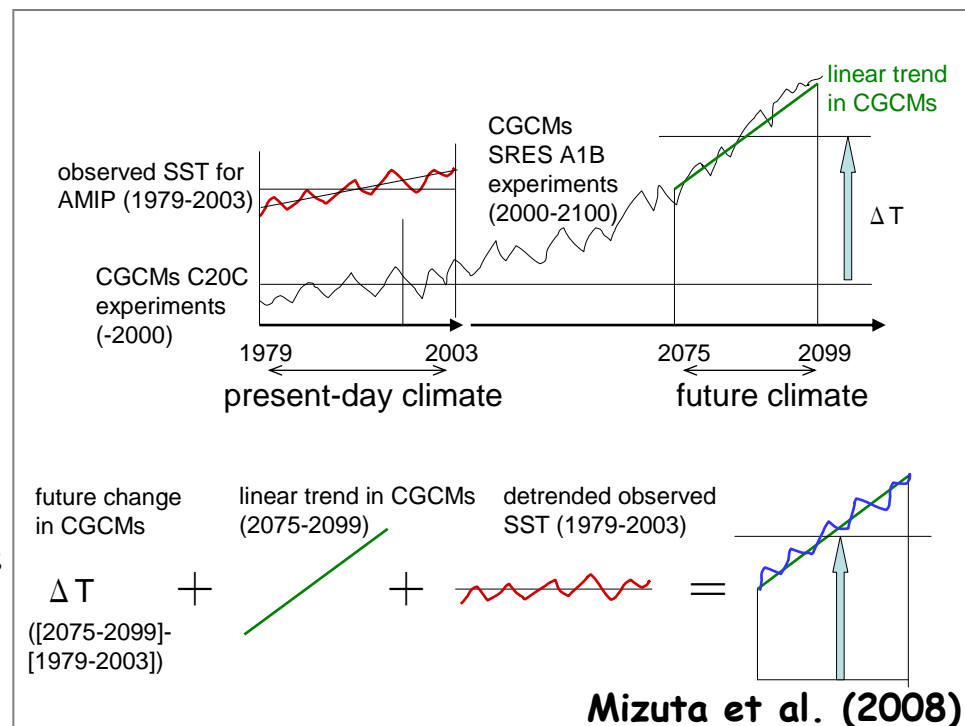


Typhoon **60km** vs **20km**

Kakushin Team-Extremes Time-Slice

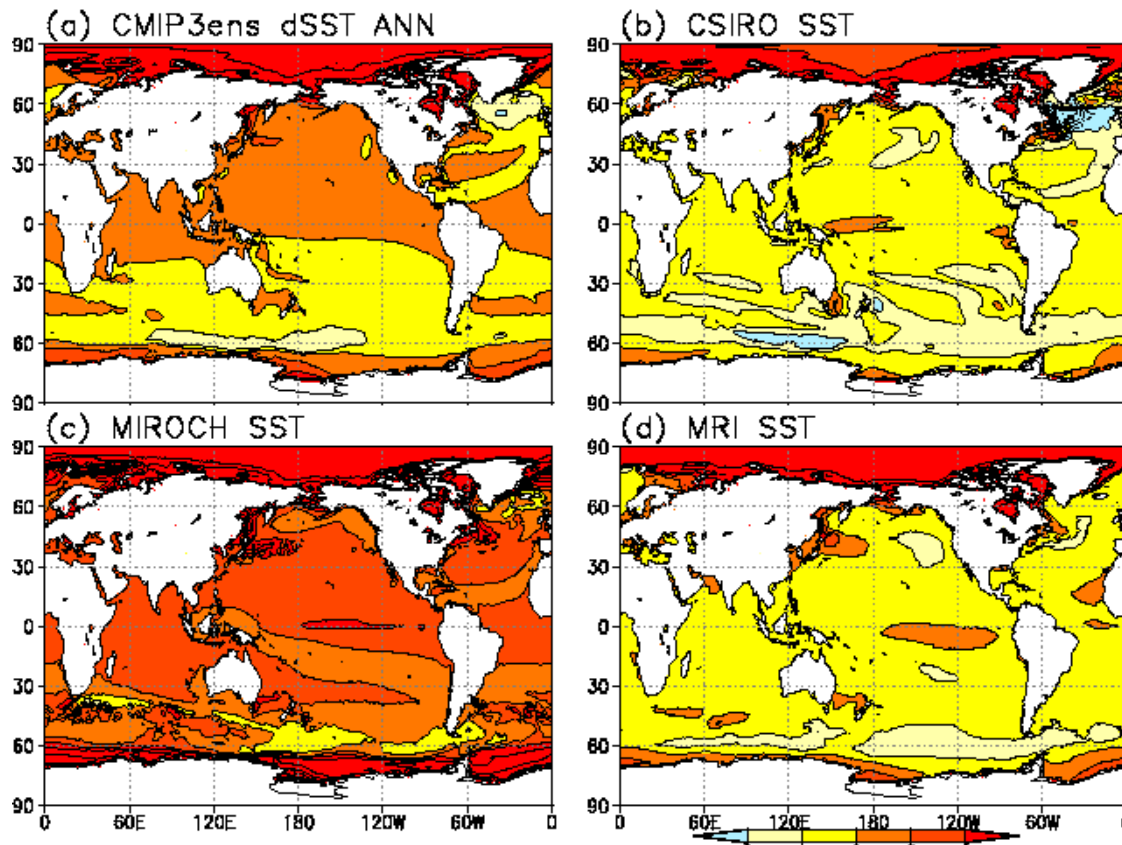


- Present-day (1979-2003)
 - the observed sea surface temperature (SST) and sea-ice concentration
- Near Future (2015-2039) and Future (2075-2099)
 - the SST and sea-ice anomalies of the CMIP3 multi-model ensemble mean are added to the observations, retaining the present interannual variability



Use of one-member 20-km AGCM run and ensemble runs with 60-km AGCM

period	SST	initial condition 0	initial condition 1	initial condition 2
Present: 1979-2003	observation	HP0A	HP0A_m01	HP0A_m02
Future: 2075-2099	CMIP3 average	HF0A	HF0A_m01	HF0A_m02
	MRI-CGCM2.3.2	HF0A_mri	HF0A_mri_m01	HF0A_mri_m02
	MIROC_hires	HF0A_miroch	HF0A_miroch_m01	HF0A_miroch_m02
	CSIRO	HF0A_csiro	HF0A_csiro_m01	HF0A_csiro_m02



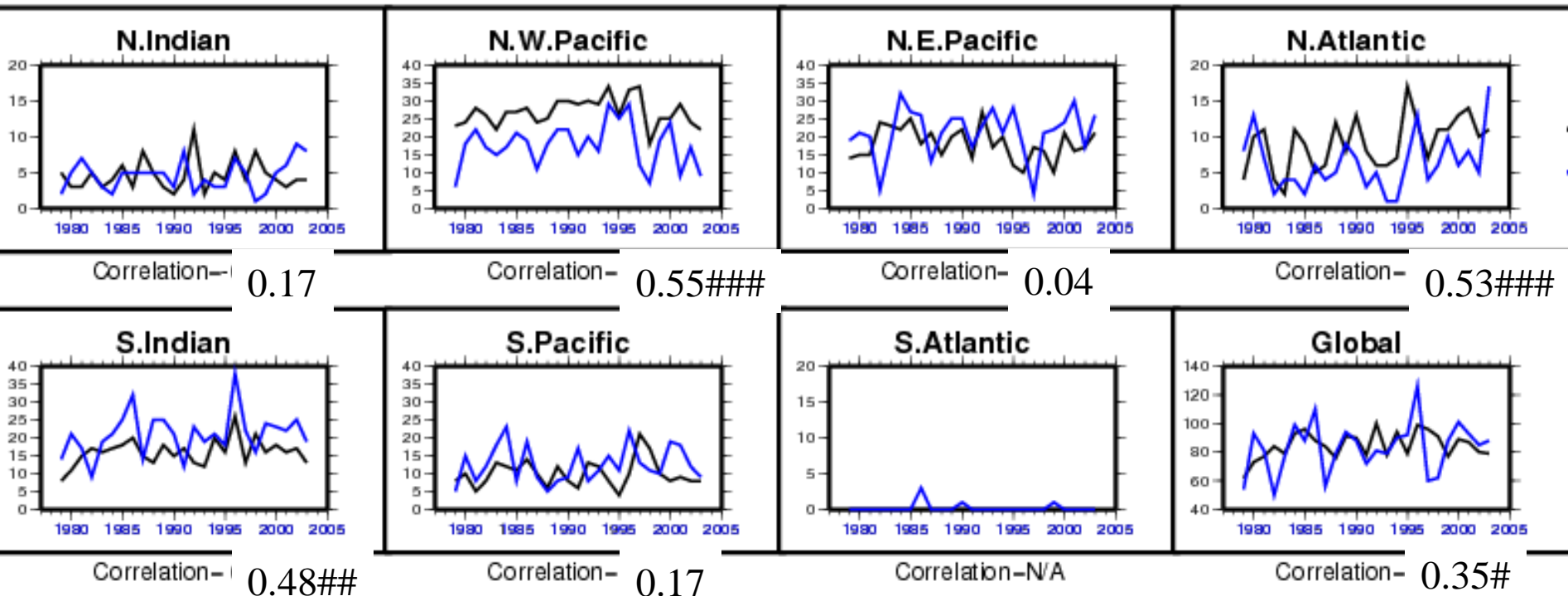
- 20-km:
1 realization
- 60-km:
4 SST anomalies
3 initial conditions

Tropical cyclones

Validation: Inter-annual variation of TC frequency

— Observation

— 20-km AGCM (AMIP run 1979-2003)



There is a skill for TC frequency interannual variation associated with ENSO

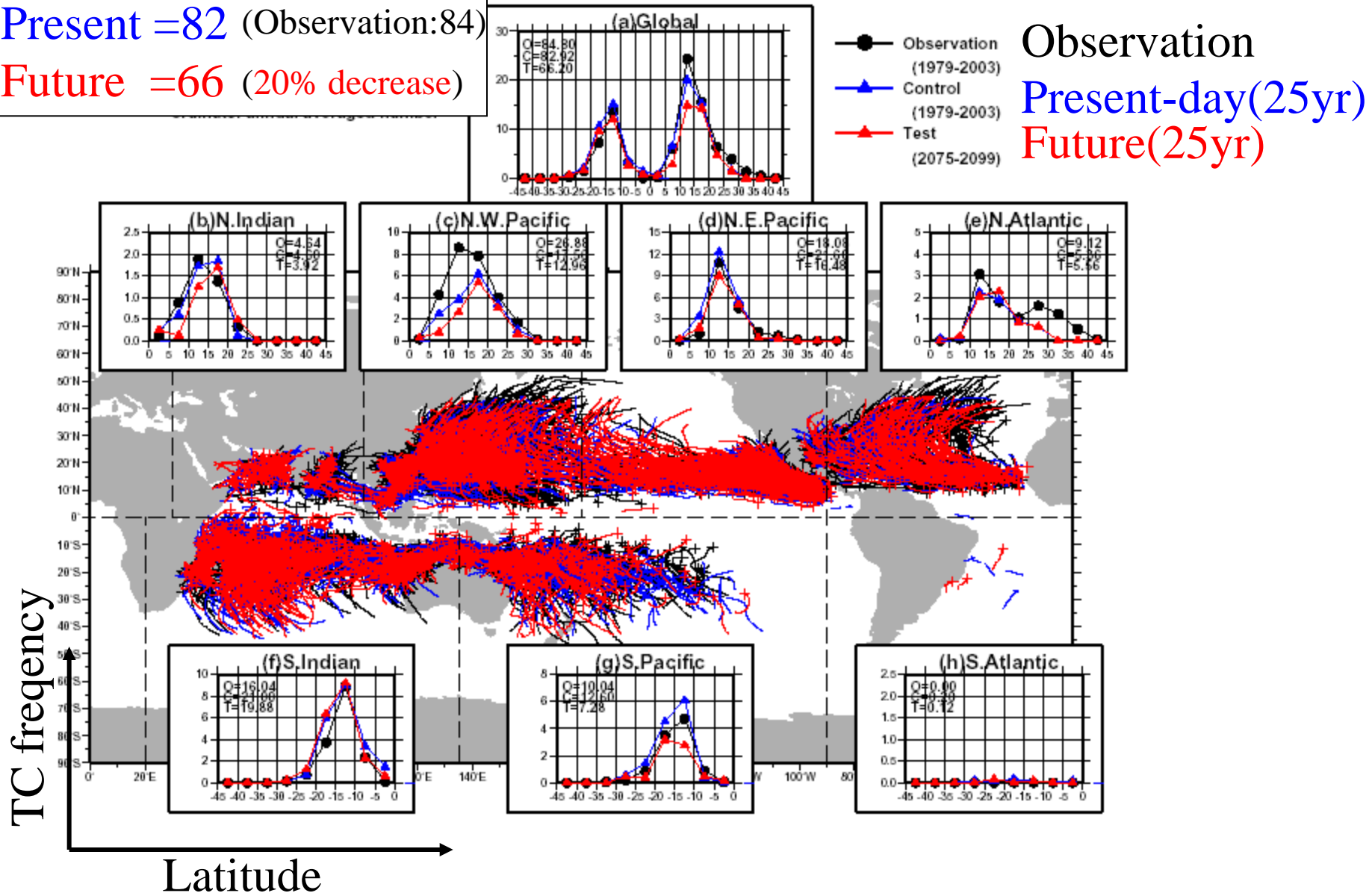
###:99% significance level
##:95% significance level
#:90% significance level

Number of TC Generated in Each Latitude

Annual global average

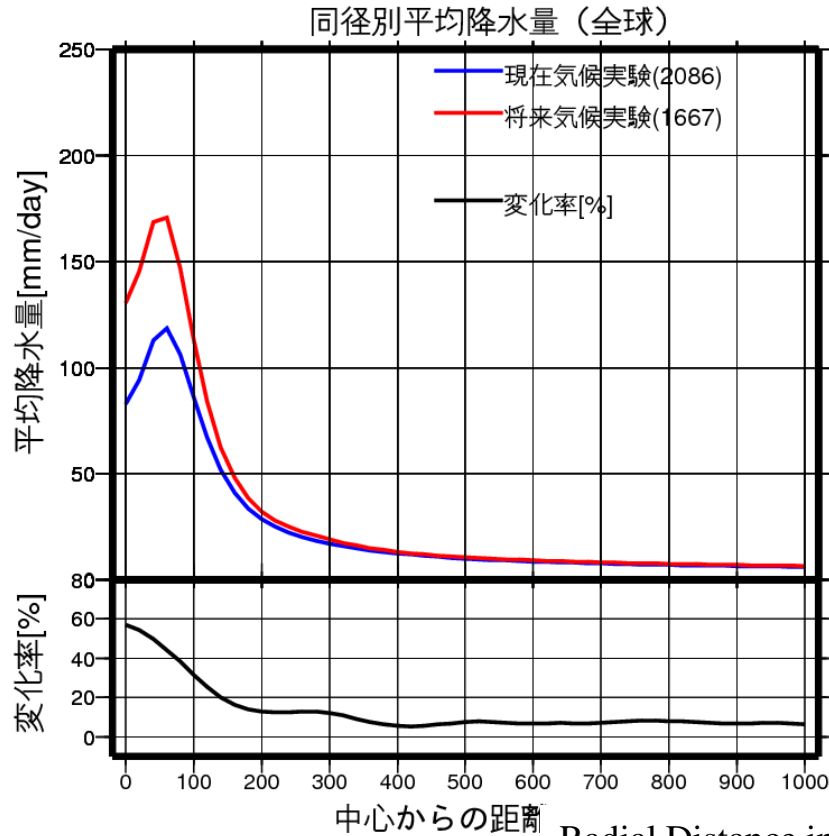
Present = 82 (Observation: 84)

Future = 66 (20% decrease)

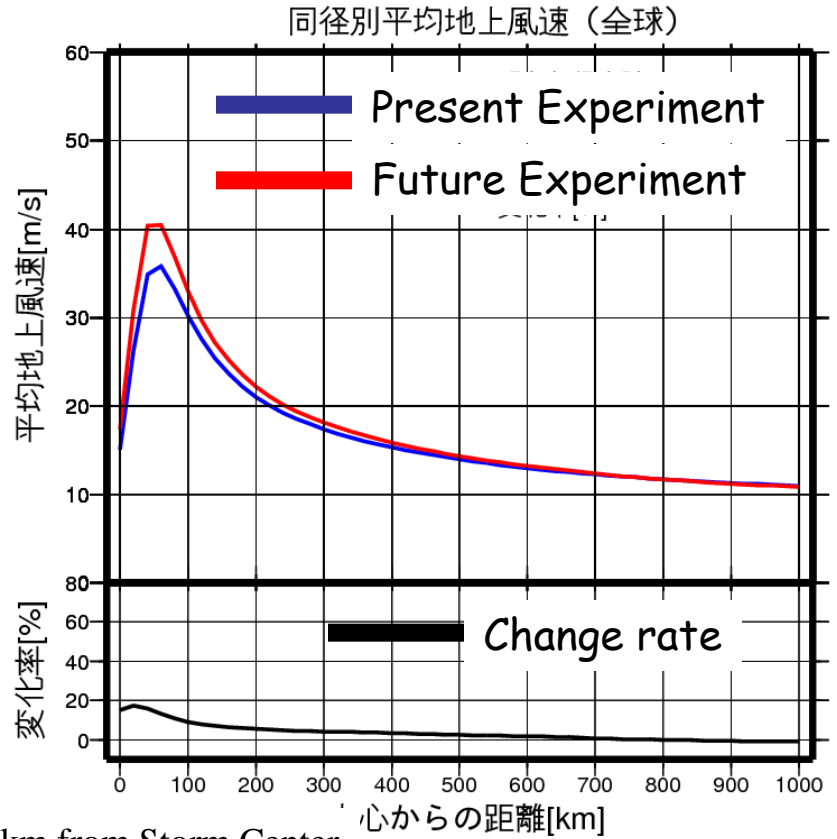


Radial Profile Change around TC

Precipitation



Surface Wind

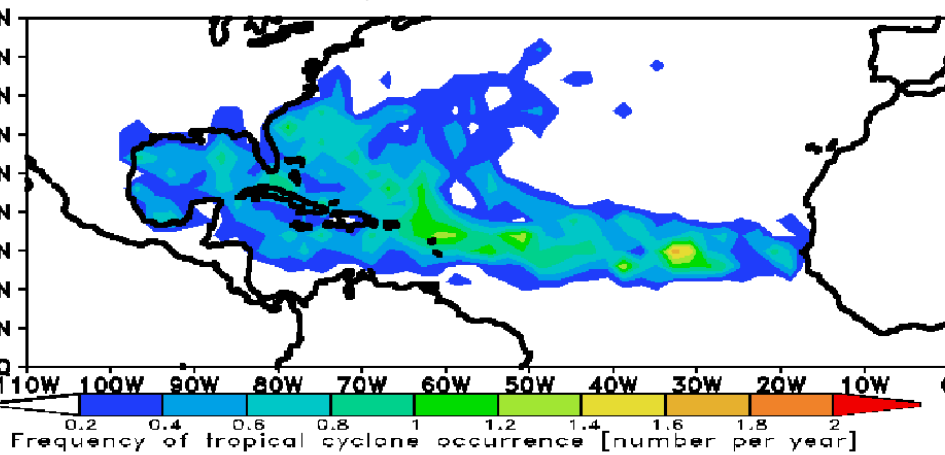


- Large changes occur near inner-core region, 40-60% for precipitation and 15-20% for surface wind.
- A surface wind speed increase of more than 4% can be seen up to 500 km from storm center.

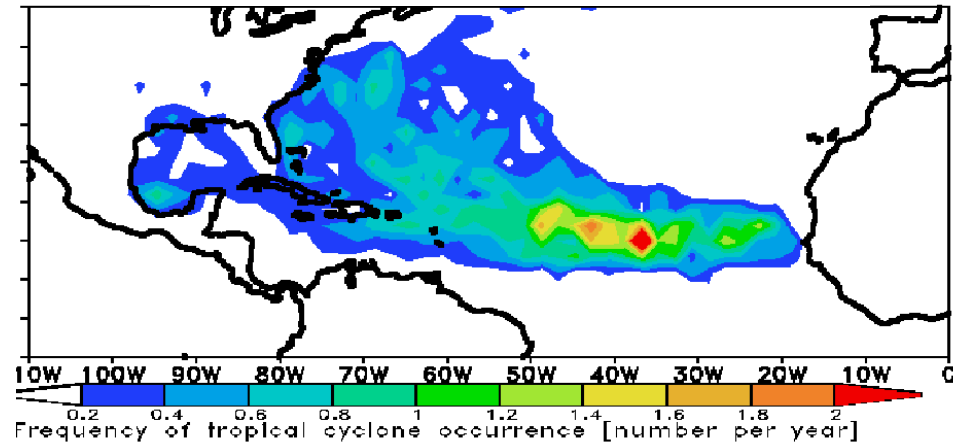
TC track changes

Future change in frequency of TC occurrence

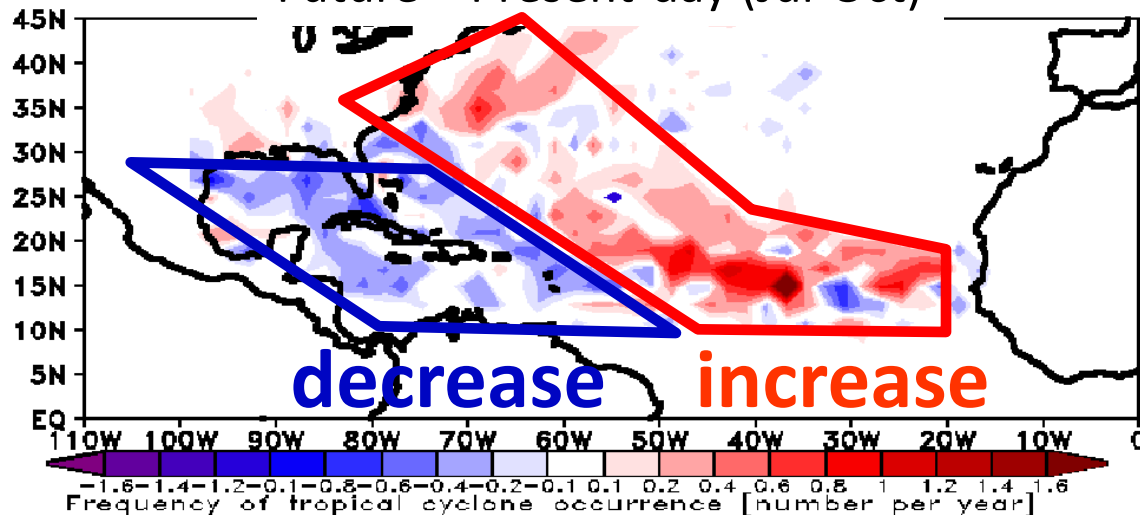
Present-day (1979-2003, Jul-Oct)



Future (2075-2000, Jul-Oct)



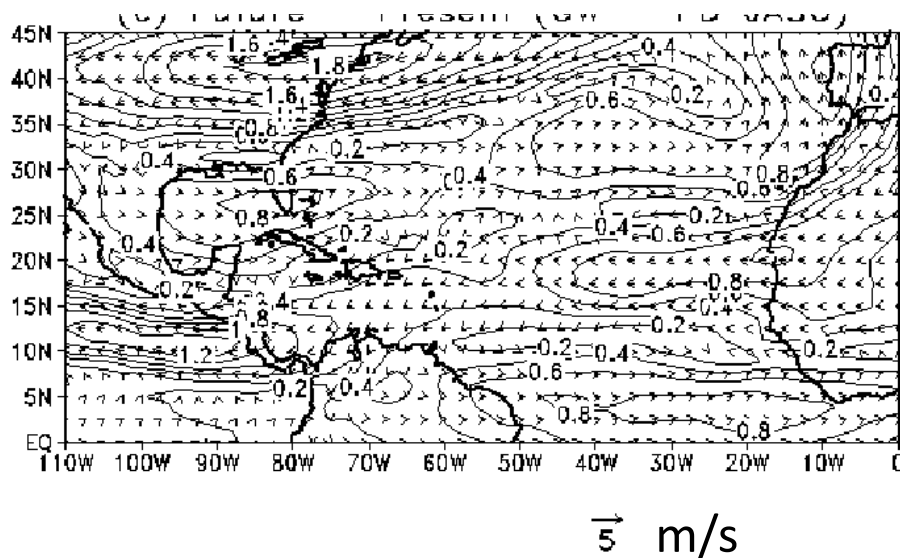
Future – Present-day (Jul-Oct)



The TC frequency will decrease in the western North Atlantic and increase in the eastern North Atlantic

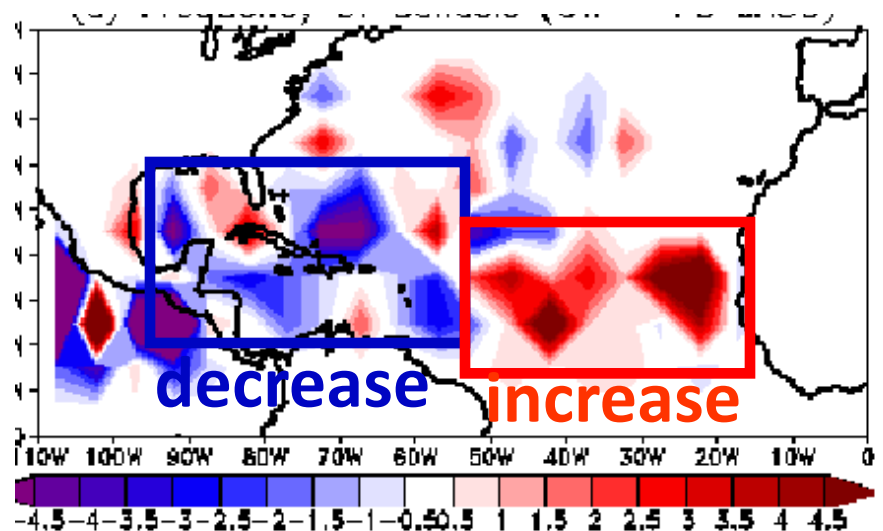
Reason for track change

Change in steering flow (Jul-Oct)



only a small difference in steering flows

Change in frequency of genesis (Jul-Oct)



Track changes are caused by alternation in TC genesis locations rather than in TC steering flows

Future change in Genesis Potential Index

$$GPI = \left| 10^5 \eta \right|^{3/2} \left(\frac{R}{50} \right)^3 \left(\frac{V_{pot}}{70} \right)^3 \left(1 + 0.1 V_{shear} \right)^2 \left(\frac{-\omega + 0.1}{0.1} \right)$$

Absolute
vorticity
at 850 hPa

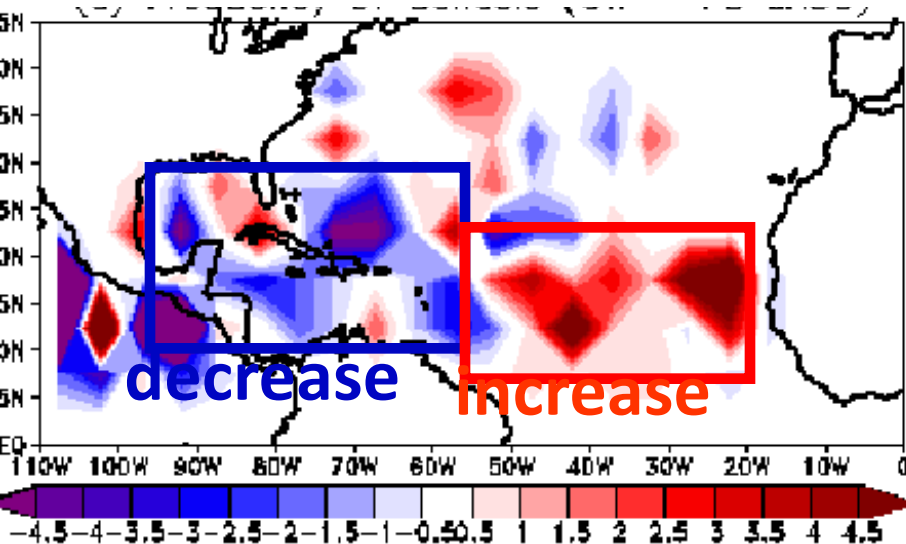
Relative
humidity
at 600 hPa

Maximum
potential
intensity

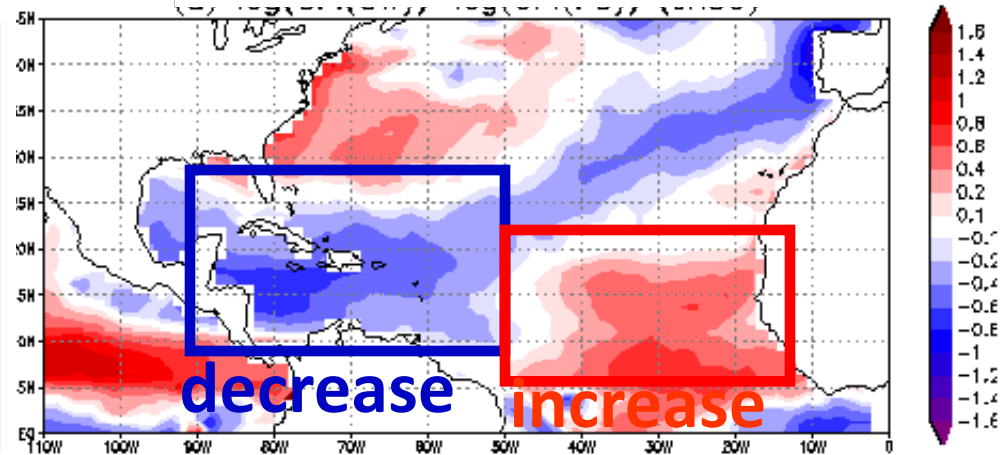
Vertical wind
shear between
850 hPa and 200 hPa

Vertical velocity
at 500 hPa

Change in frequency of genesis (Jul-Oct)



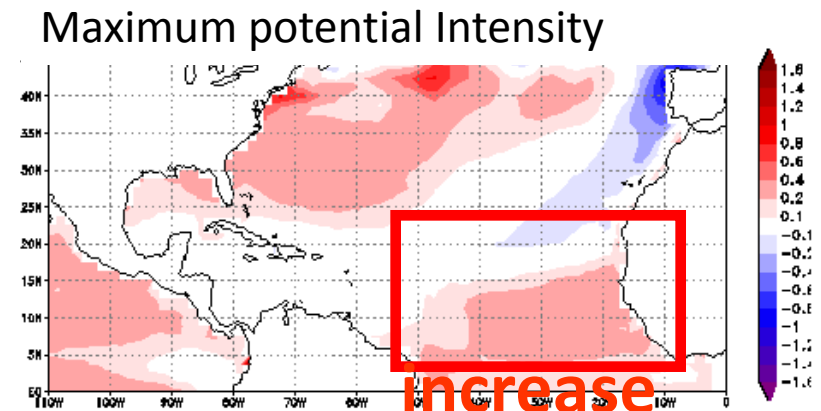
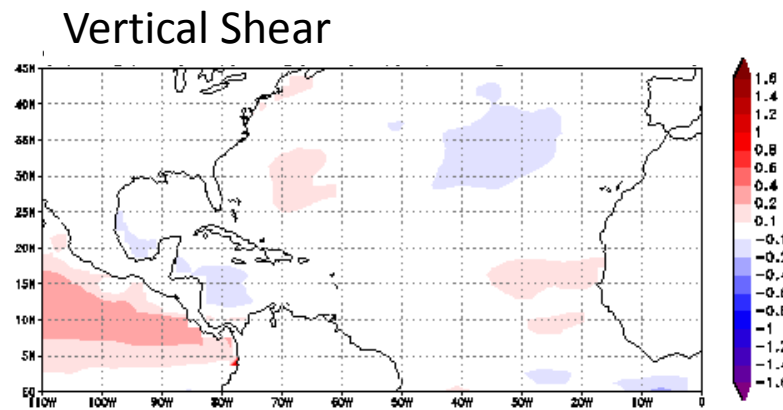
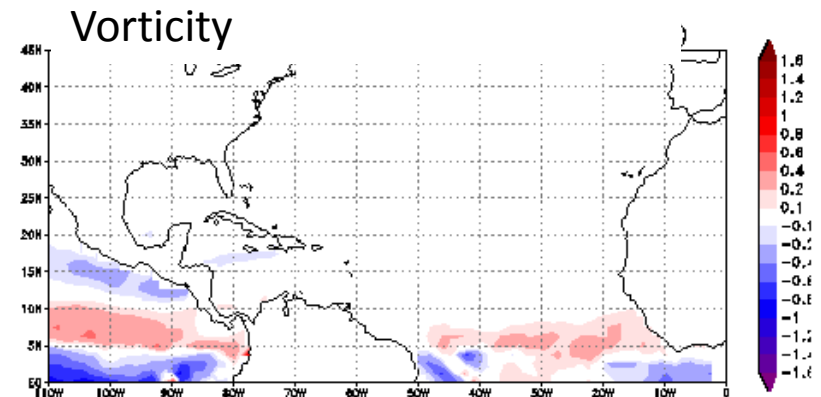
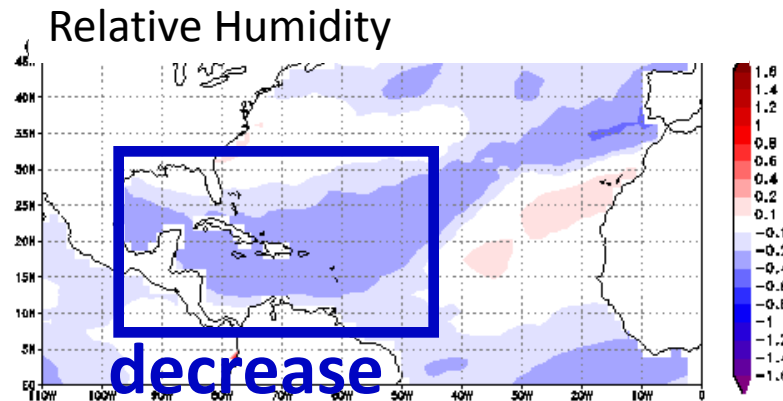
Change in GPI (Jul-Oct)



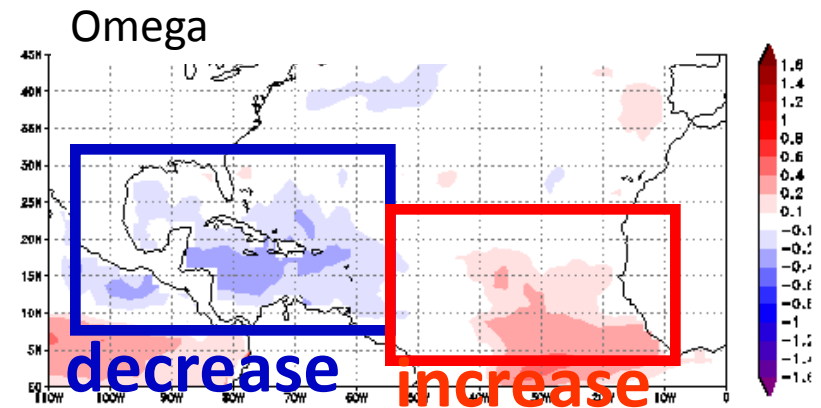
The GPI change correctly depicts the change in frequency of genesis

The GPI can be used to determine which of the GPI elements contribute most to its future change

Influential factor to the change in GPI



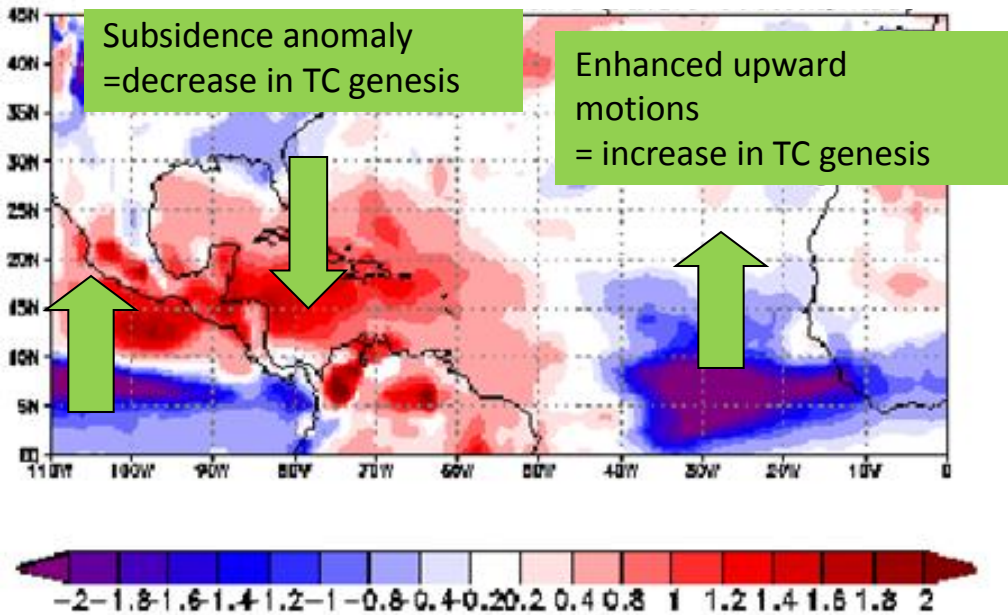
Changes in the maximum potential intensity and omega terms make the dominant contribution to the increase in the GPI within the eastern North Atlantic, whereas the relative humidity and omega terms make the largest contribution to the decrease in the GPI within the western North Atlantic



See poster by H. Murakami

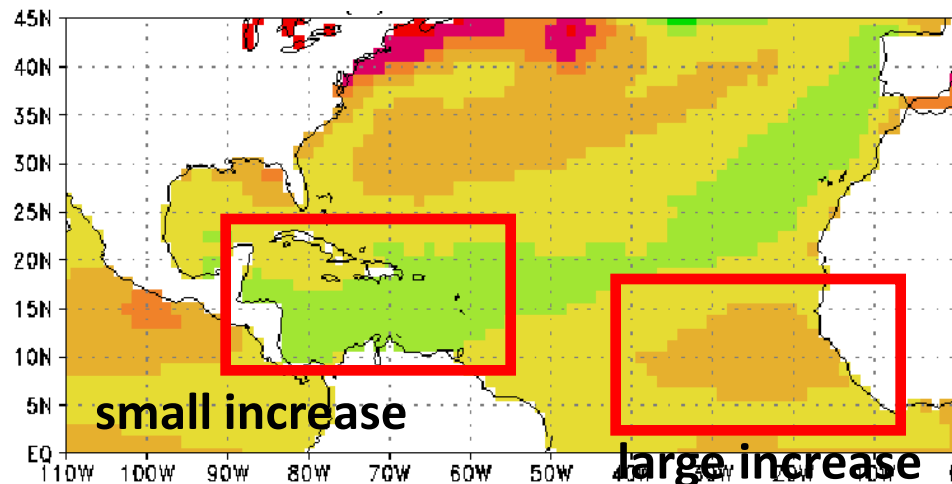
Reason for TC genesis change

Change in Omega (Jul-Oct)



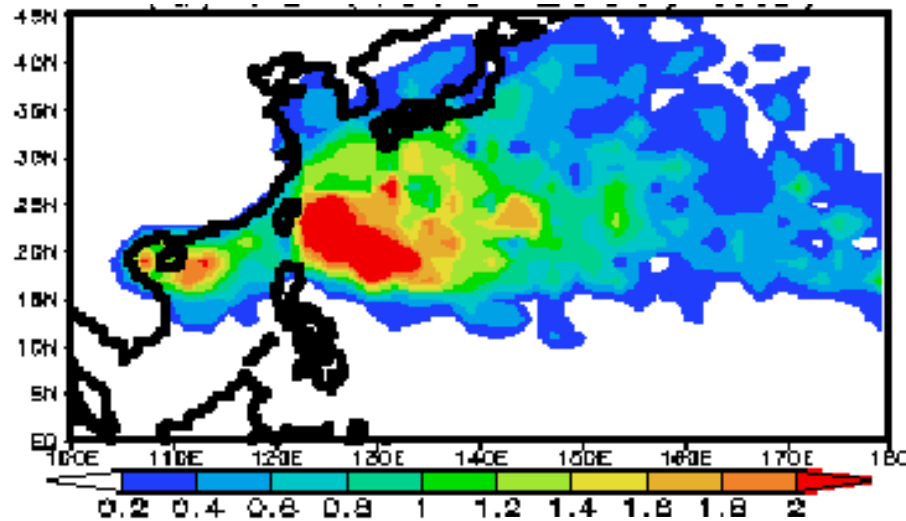
The reduction of genesis in the west North Atlantic is attributed to decreases in relative humidity and ascending motion caused by an enhanced large scale subsidence, whereas the increase of genesis in the southeast North Atlantic arises from increasing upward motion and convective available potential energy due to local ocean surface warming

Change in prescribed SST (Jul-Oct)

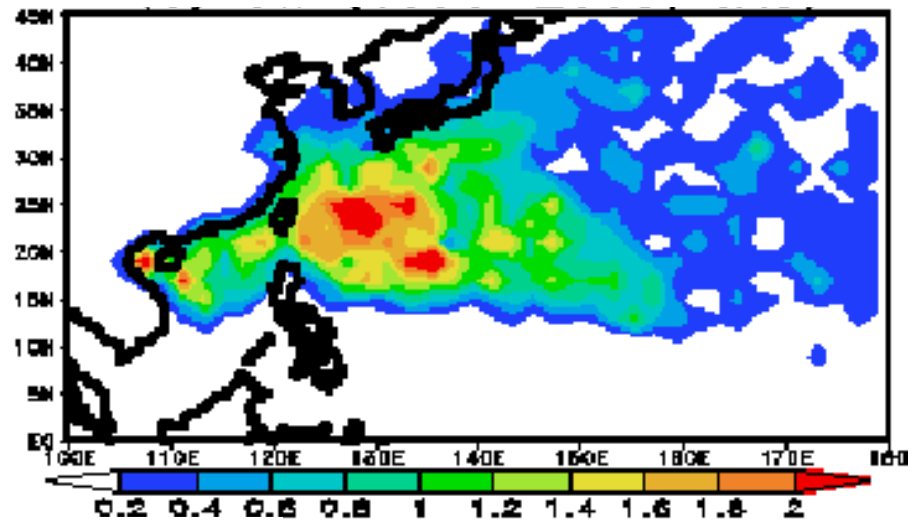


Future change in frequency of TC occurrence

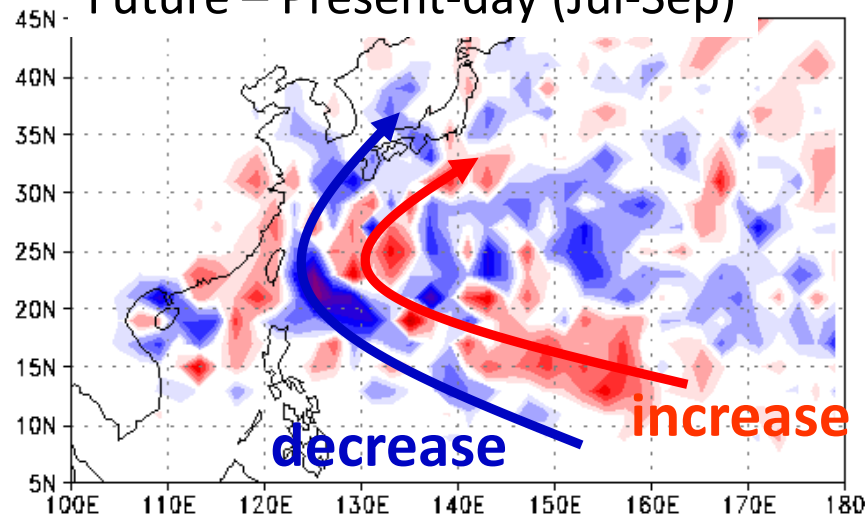
Present-day (1979-2003, Jul-Sep)



Future (2075-2000, Jul-Sep)



Future – Present-day (Jul-Sep)



Frequency of TC landfall over Japan and Korea may decrease

**Future change in tropical cyclone
frequency with various resolutions
and SST settings**

Future change in TC frequency

The changes in tropical cyclone frequency as projected by 20-km mesh and 60-km mesh global atmospheric model experiments. The changes are shown in terms of the ratio of future frequency to present frequency. Statistically significant increase (decrease) at 95% confidence level by two-sided t-test is indicated by red (blue) color.

Experiments	Resolution	Δ SST	Integration	Ratio (%) of TC frequency Future/Present								
				Global	NH	SH	N Indian	NW Pacific	NE Pacific	N Atlantic	S Indian	S Pacific
A0	TL959, 20km	MRI CGCM 2.3	20yr	71	69	73	61	64	61	122	72	77
A1		MRI CGCM 2.3	20yr	75	75	75	71	71	70	123	75	73
A2		MIROC-H	10yr	73	85	58	132	128	50	82	76	11
A3		CMIP3	25yr	80	79	81	85	74	75	105	95	58
B1	TL319, 60km	MRI CGCM 2.3	25yr	80	79	83	89	66	69	150	78	92
B2		MIROC-H	25yr	94	100	84	179	164	58	106	110	31
B3		CMIP3	25yr	79	81	76	133	86	67	104	79	64
B4		CSIRO	25yr	78	71	89	93	113	51	63	78	110
C3	TL159, 120km	CMIP3	25yr	71	79	54	99	75	78	75	62	38

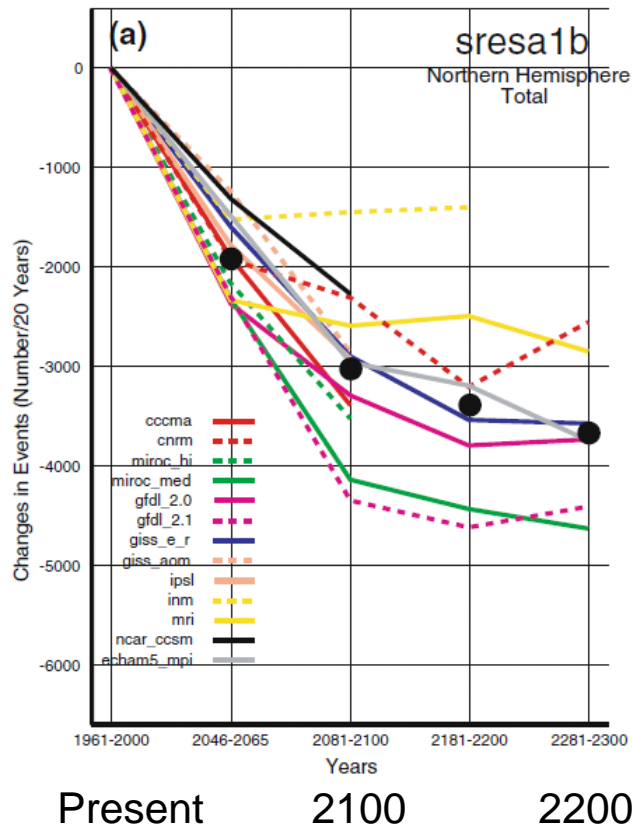
Commonly reduced regardless of difference in SST increase pattern.

Mostly same changes are achieved by the same SST setting regardless of resolution difference.

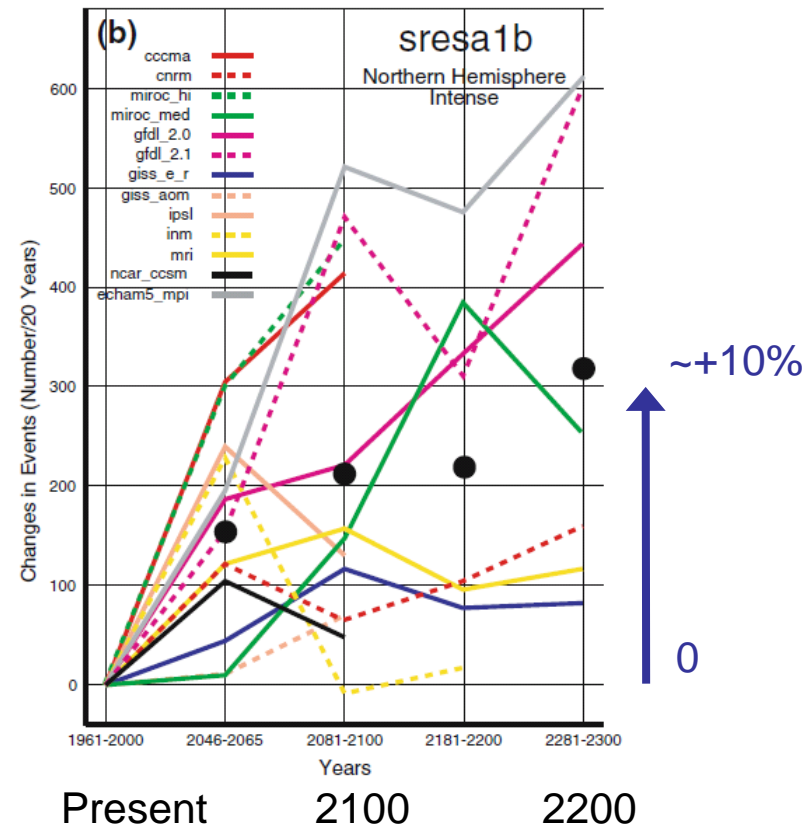
Extratropical cyclones

Extratropical Cyclones

Total cyclone number



“Strong” cyclone number (<970hPa)

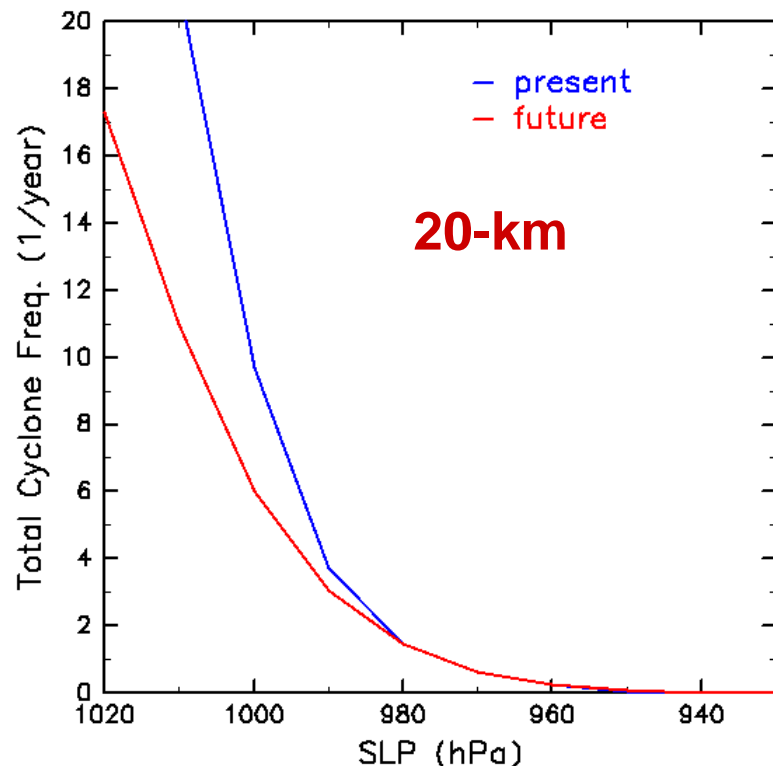


- When tracking extratropical cyclones..
 - Number of cyclones decreases
 - “Strong” cyclones increase

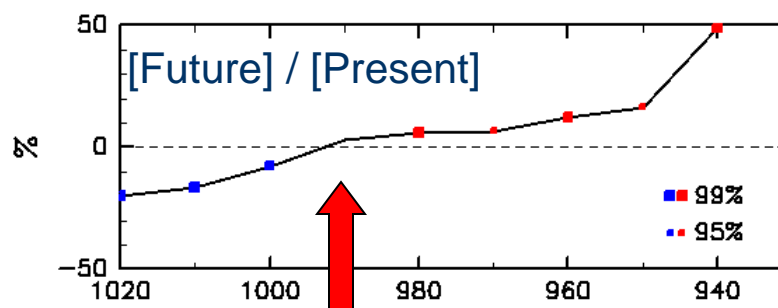
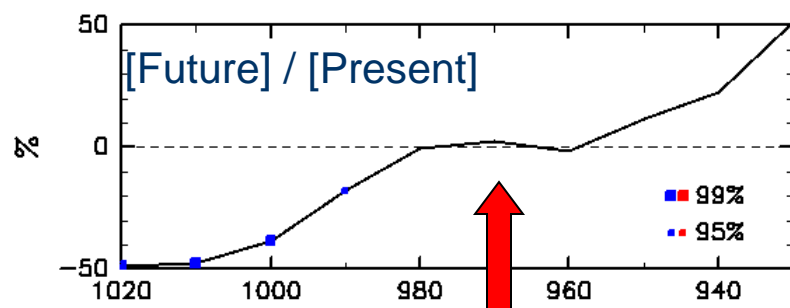
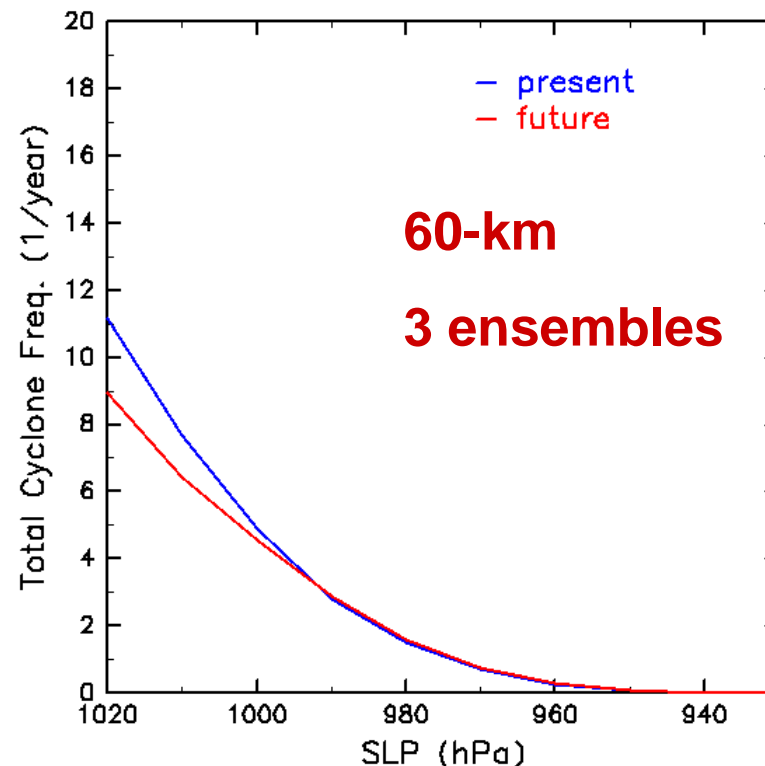
Same in high-resolution models but with different threshold for + or -

Frequency of cyclones as a function of threshold pressure

($\times 1000$) SPOA SFOA 20N-90N DJF (≥ 12 h)

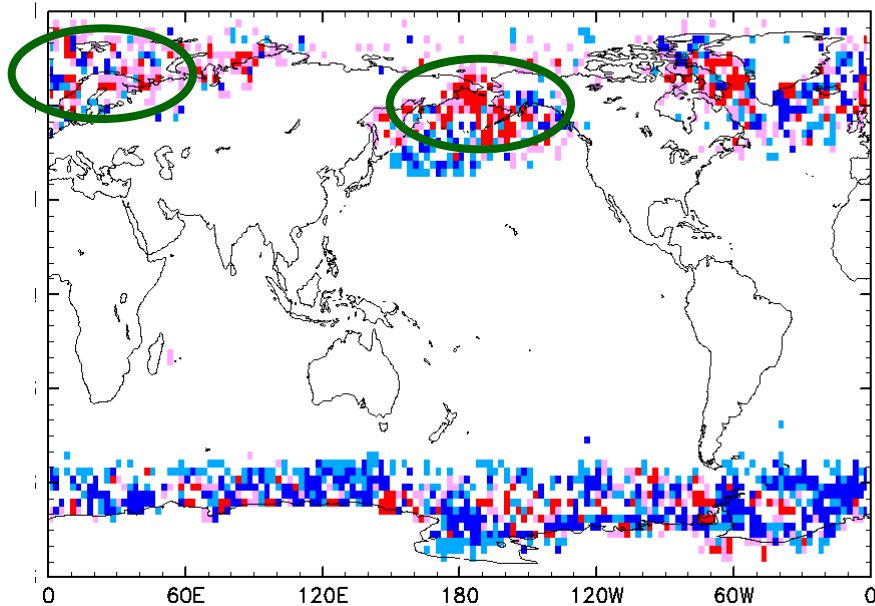


($\times 1000$) HPx3 HFx3 20N-90N DJF (≥ 12 h)



Cyclone Frequency ($\geq 12\text{h}$, $\leq 980\text{hPa}$)

DJF HFx3-HPx3

**Strong Cyclone Frequency**

[Future] - [Present]

Strong cyclones increase in the downstream of the storm tracks

Cyclone Growth Rate

[Future] - [Present]

Growth Rate = temporal change of

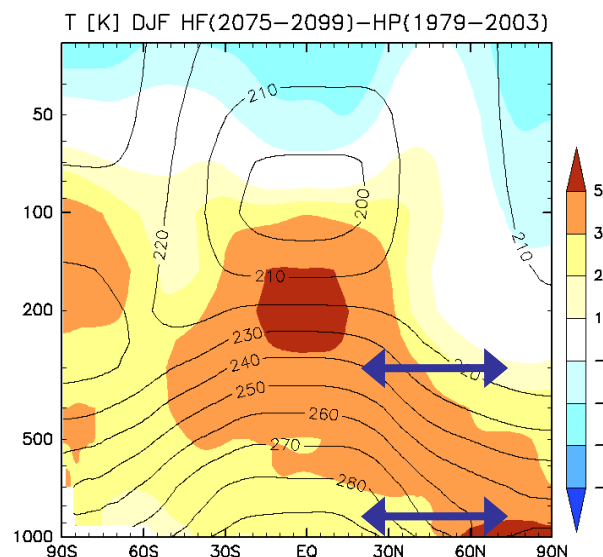
$$\frac{\partial \text{SLP}}{\partial t} = -(\text{SLP}_{t+} - \text{SLP}_{t-})$$

Increase in the upstream of strong cyclone increases

Baroclinicity (maximum Eady growth rate; Lindzen and Farrell 1980)

$$\sigma_{BI} = 0.31 g N^{-1} T^{-1} \nabla T$$

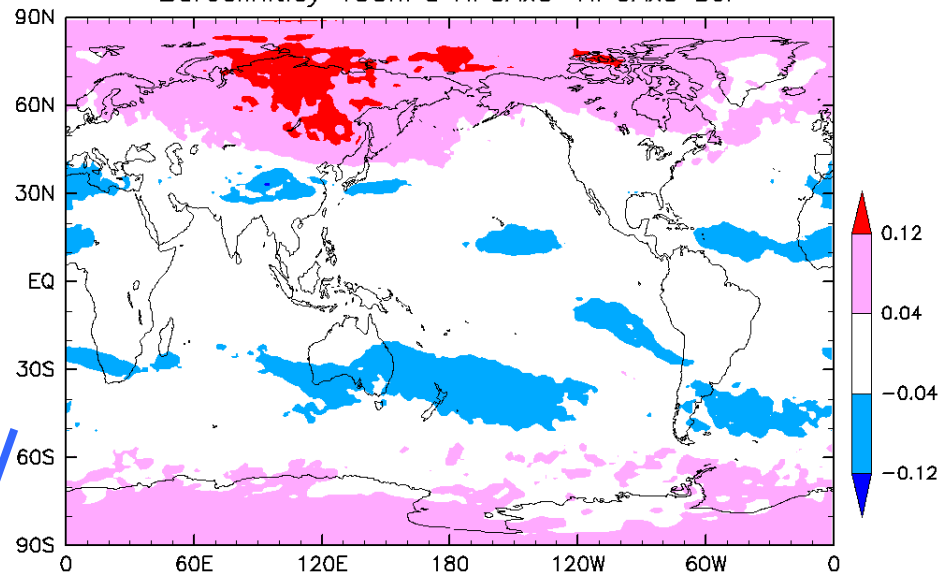
Increase at higher levels of polar region



MRI AGCM

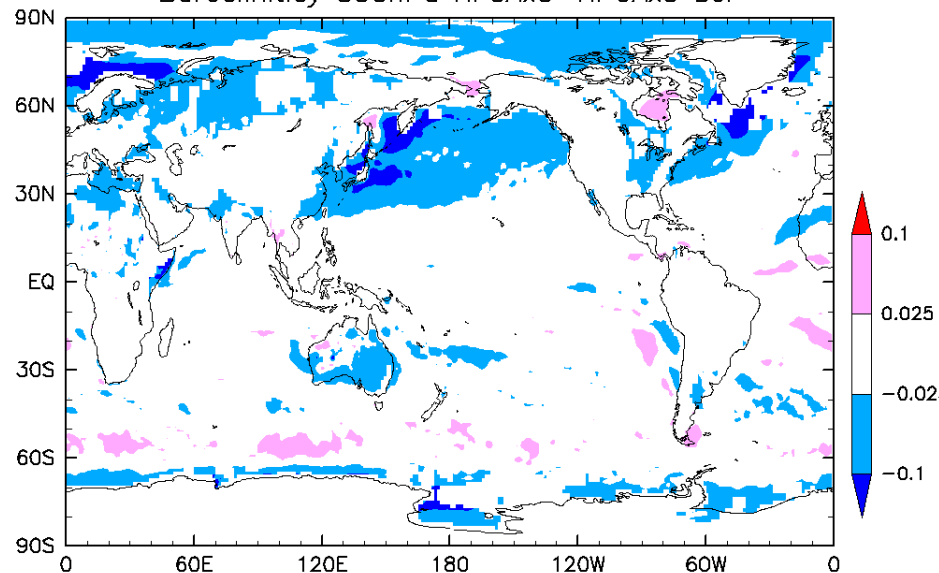
400 hPa

Baroclinicity 400hPa HF0Ax3–HP0Ax3 DJF



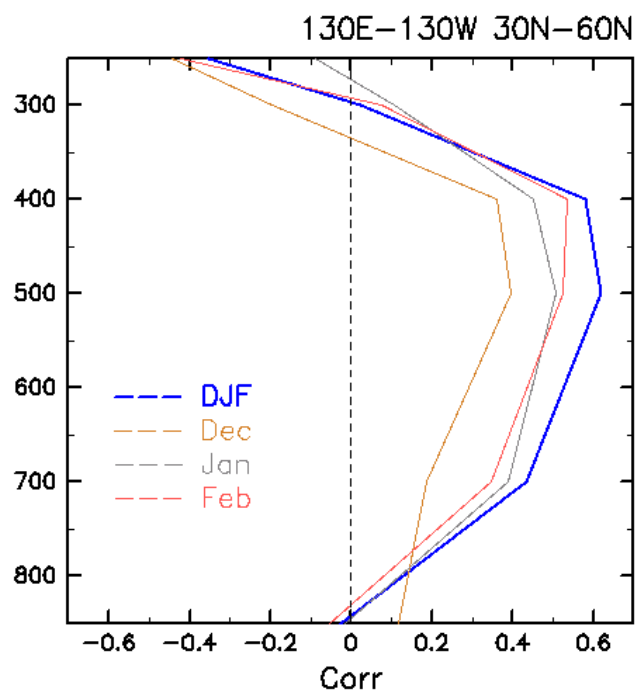
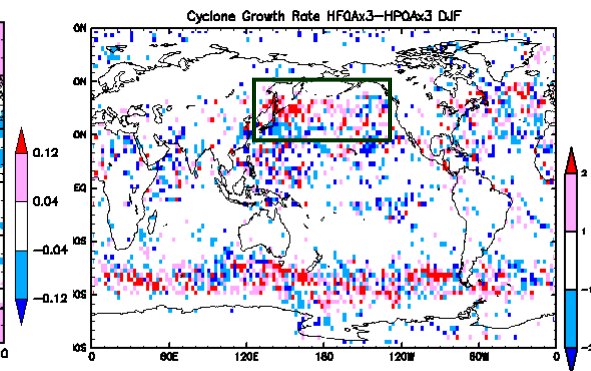
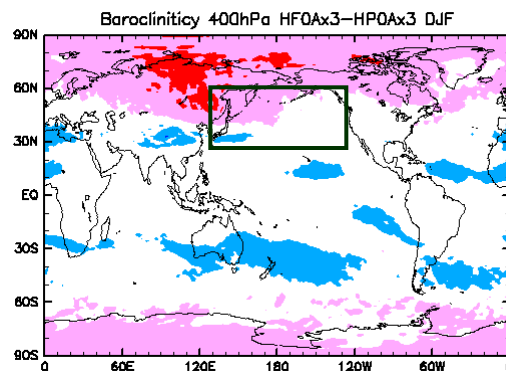
850 hPa

Baroclinicity 850hPa HF0Ax3–HP0Ax3 DJF

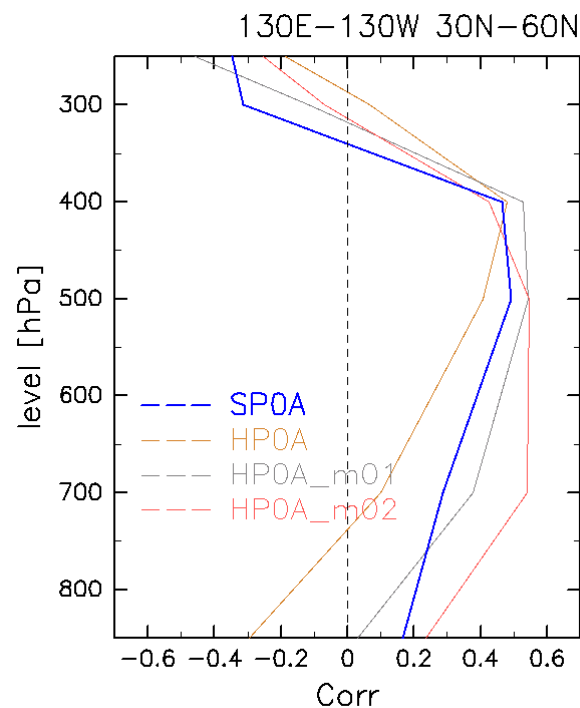


Mizuta et al. (2009)

Pattern correlation
between the cyclone
growth rate change and
baroclinicity change
(Pacific region)



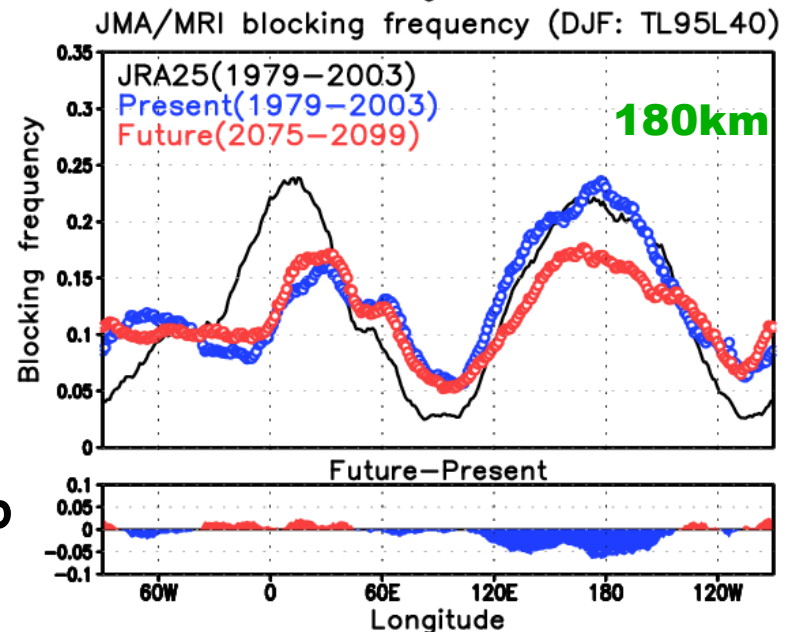
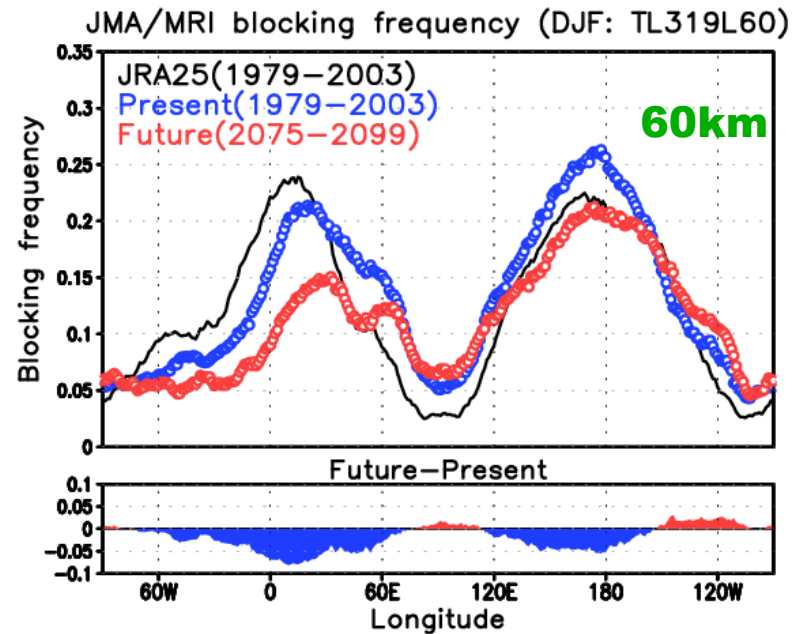
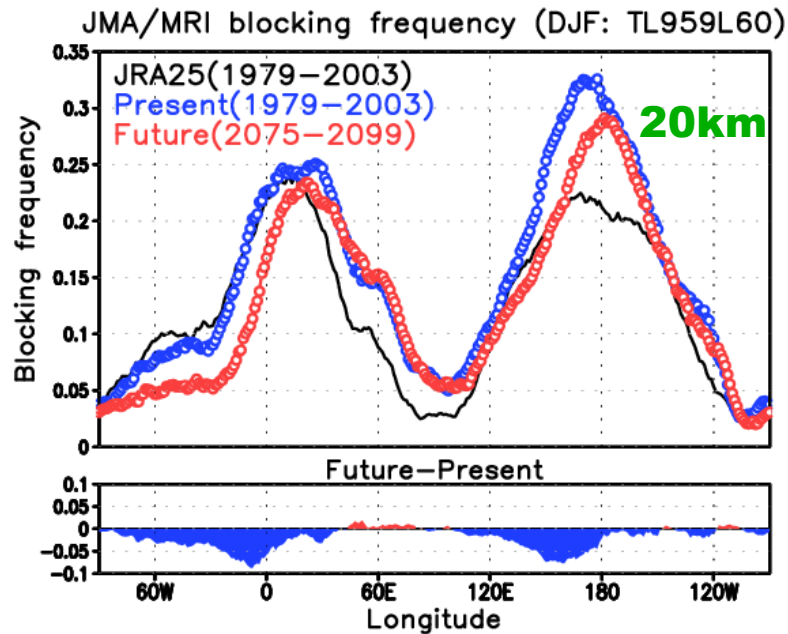
Each month of 3-member ensembles



DJF of each member

NH winter blockings

Northern Hemisphere wintertime blocking

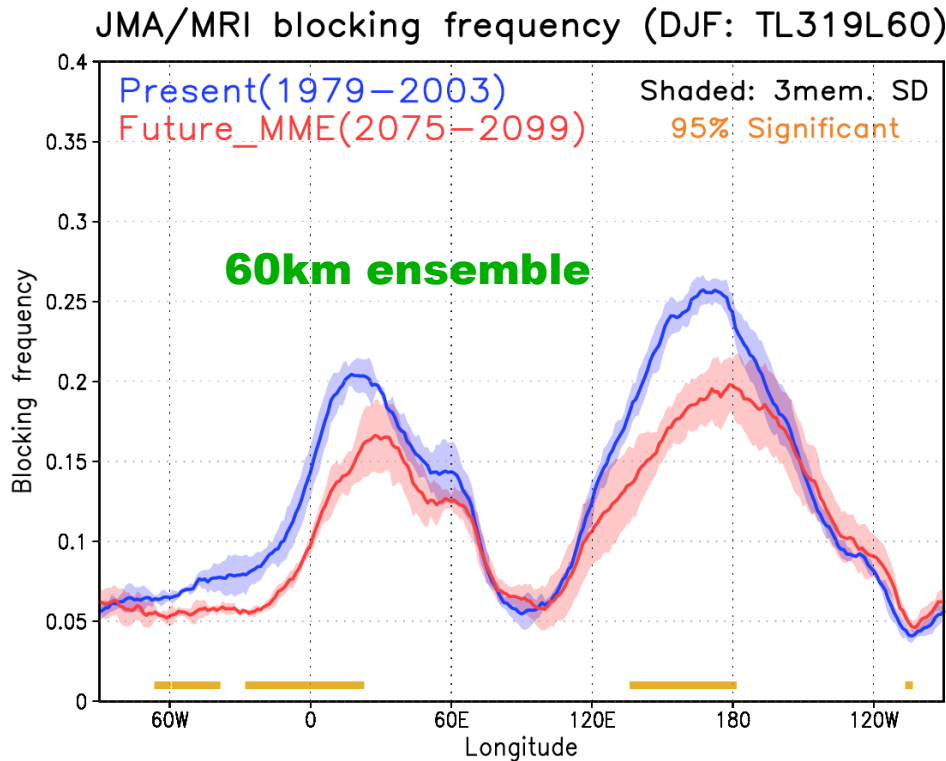


The higher the resolution, the more accurate the simulated Euro-Atlantic blocking frequency

Frequencies of Euro-Atlantic and Pacific blockings are projected to decrease

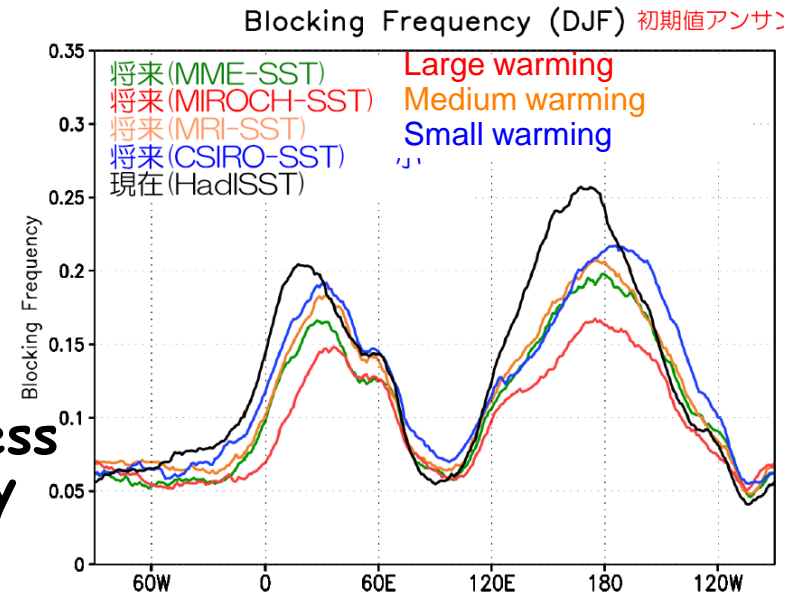
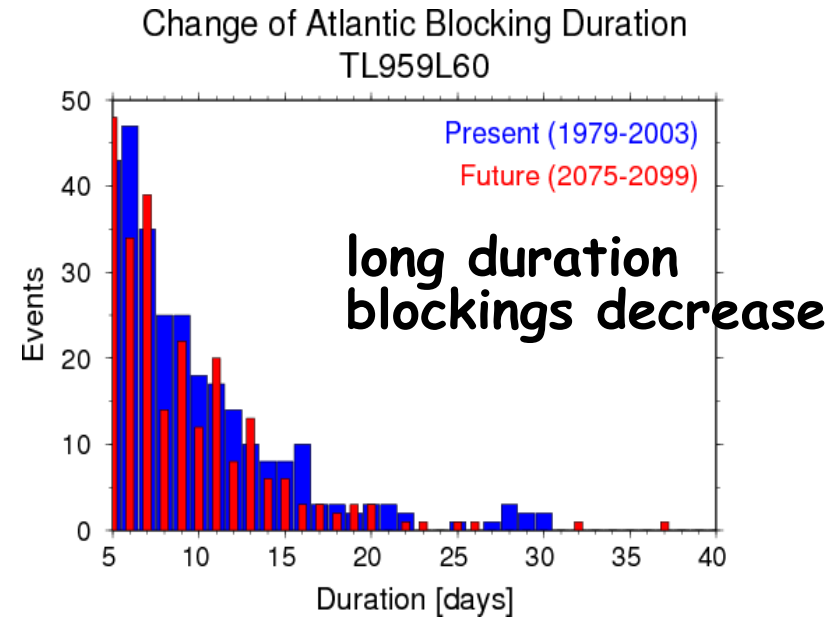
Uncertainty in future projections of blocking

60km ensemble simulations with different SSTs



frequencies of Euro-Atlantic and Pacific blockings are projected to decrease significantly.

The larger the warming is, the less blocking frequency

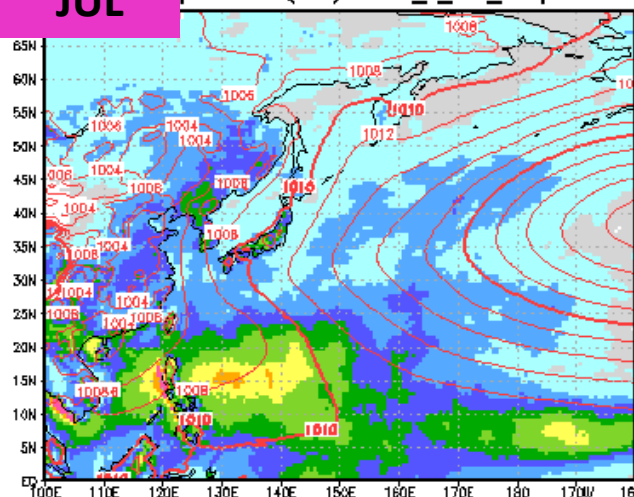


Future plan

- Next 2 years: use of improved model for TC and precipitation
AOGCM (ESM): CMIP5 runs with TL159 AGCM + 1x0.5 OGCM
with TL95 aerosol model + T42 chemistry model
AGCM: TL959 (20-km) AGCM time-slice

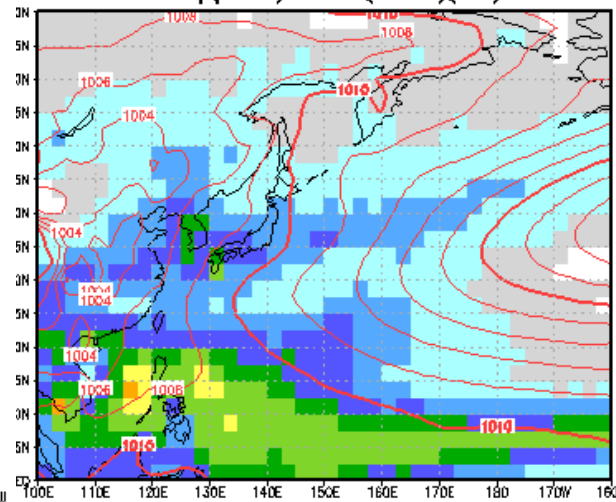
New Cumulus + Ocean Skin

JUL Precip & SLP (JUL) HPAI_Y_skn_omip



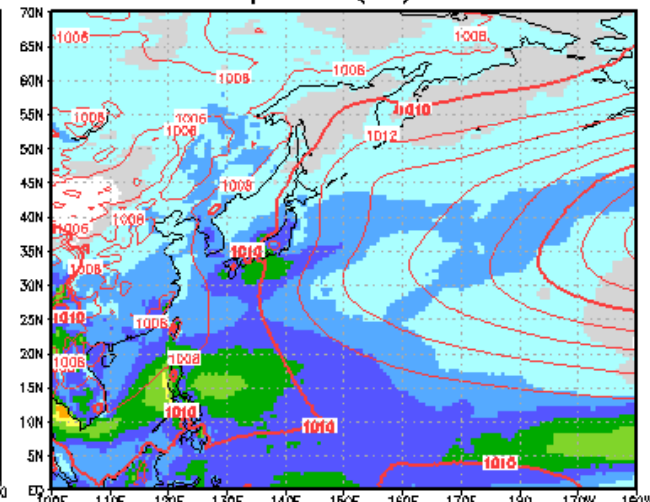
OBS

(ReanalysisCMAP)
Precip(CMAP) & SLP(ERA40)(JUL)



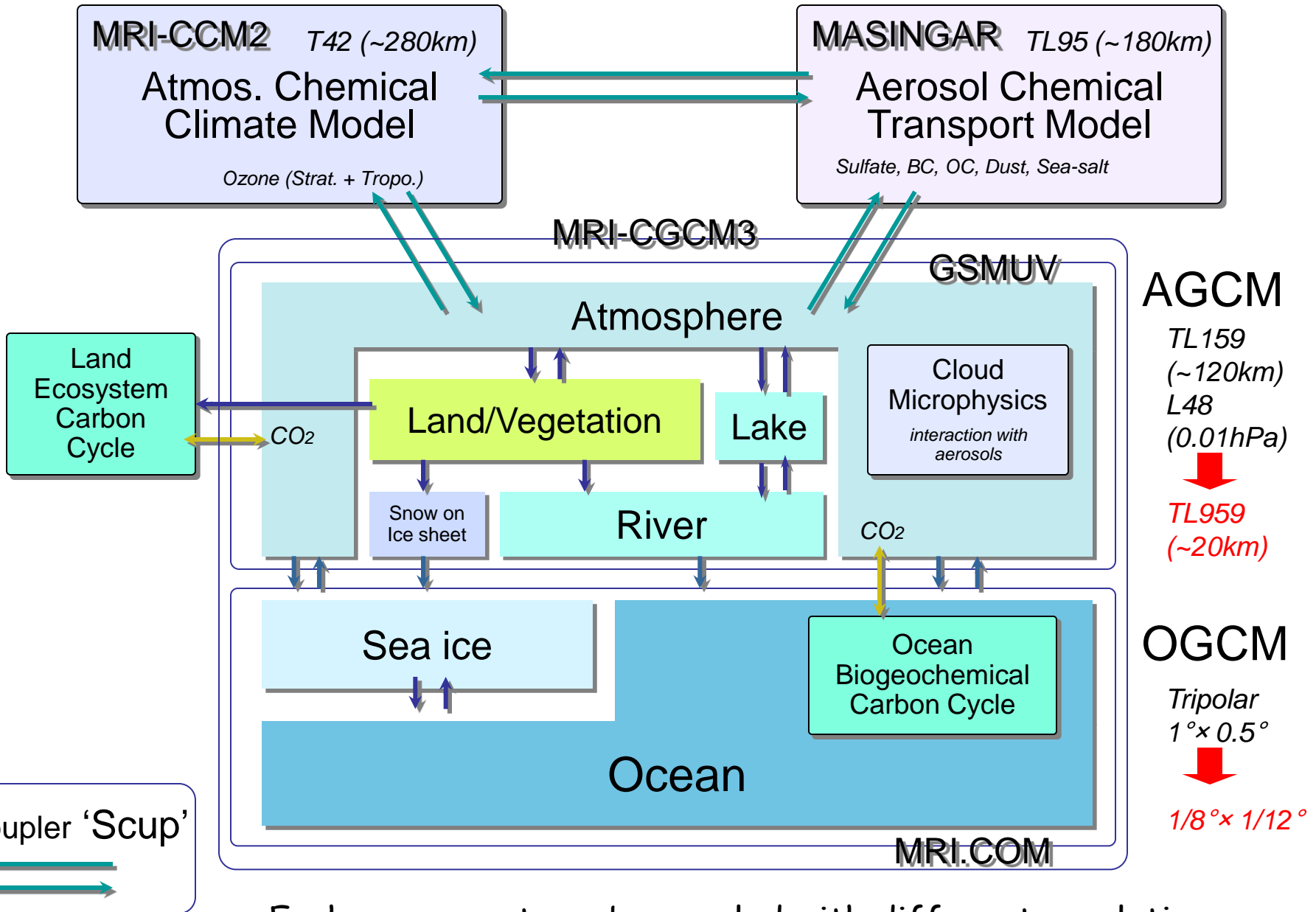
Current Model

Precip & SLP (JUL) HP0A



- After that: full AOGCM + ESM
Atm: 20-km, Ocn: 1/8x1/12,
plus lower resolution aerosol and chemistry model

MRI Earth System Model: CMIP5 and after



Each component can be coupled with different resolutions